Coupled Feedbacks on the Northern Hemisphere Midlatitude Jet Response to 4xCO_2: Changes in Stratospheric Ozone and the Atlantic Meridional Overturning Circulation

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ABSTRACT: Ozone, and its response to anthropogenic forcings, provide an important pathway for the coupling between atmospheric composition and climate. This applies to stratospheric ozone as well as ozone in the troposphere; in addition to stratospheric ozone's radiative impacts, recent studies have shown that changes in the ozone layer due to $4xCO_2$ have a considerable impact on the Northern Hemisphere (NH) tropospheric circulation, inducing an equatorward shift of the North Atlantic jet during boreal winter. Here we show that this equatorward jet shift induces a more rapid weakening of the Atlantic Meridional Overturning Circulation (AMOC), resulting in a poleward shift of the jet on longer timescales. As such, coupled feedbacks from both stratospheric ozone and the AMOC result in a two-timescale response of the NH midlatitude jet to abrupt $4xCO_2$ forcing: a "fast" response (5-20 years) during which the North Atlantic jet shifts equatorward and a "long" response (~100-150 years) during which the jet shifts poleward. The latter is driven by a weakening of the AMOC that develops in response to weaker surface zonal winds, that result in reduced heat fluxes out of the subpolar gyre, reducing North Atlantic Deep Water formation. Our results suggest that stratospheric ozone changes in the tropical lower stratosphere can have a surprisingly powerful effect on the AMOC, independent of other aspects of climate change.

1. Introduction

There is large uncertainty in the atmospheric circulation response to increasing greenhouse gases (e.g., Shepherd (2014)). Although models generally predict a poleward shift of the westerly jet, the magnitude of this shift is highly uncertain (e.g., Vallis et al. (2015); Grise and Polvani (2014)) as are its underlying drivers (Shaw (2019)). This is especially true in the Northern Hemisphere (NH), where there are opposing thermodynamic influences, i.e. opposite meridional temperature gradient responses at the surface versus the upper troposphere (Shaw et al. (2016)). Thus, while enhanced warming in the lower polar troposphere relative to the lower tropical troposphere (i.e., Arctic amplification) contributes to reduced meridional temperature gradients, increases in upper tropospheric tropical warming contribute to enhanced temperature gradients aloft (Butler et al. (2010); Yuval and Kaspi (2020)) and it is not clear how these competing processes affect the zonal mean jet.

Many processes have been shown to influence the response of meridional temperature gradients to increased CO_2, including polar amplification (see Smith et al. (2019) and references therein) and cloud feedbacks (e.g., Ceppi and Hartmann (2015); Voigt and Shaw (2015)). By comparison, composition feedbacks associated with the ozone response to CO_2 have been less well examined although stratospheric ozone changes have been identified as an important pathway coupling composition to climate (Isaksen et al. (2009)). In particular, the stratospheric ozone response to $4xCO_2$ consists of robust decreases in the tropical lower stratosphere (LS), increases in the tropical upper stratosphere and increases over high latitudes (Chiodo et al. (2018)). While the exact details of these changes are model dependent, especially over high latitudes, the general pattern is very consistent among models (Nowack et al. (2015), Chiodo et al. (2018), Chiodo and Polvani (2019) (hereafter CP2019)).

This pattern of reduced (increased) ozone over the tropical lower (high latitude) LS in response to $4xCO_2$ has immediate implications for temperature gradients in the stratosphere by cooling the tropics and warming high latitudes (Nowack et al. (2015); Chiodo et al. (2018)). As CP2019 showed, these changes in temperature gradients drive an anomalous equatorward shift of the midlatitude jet, not only in the Southern Hemisphere (SH), but also in the Northern Hemisphere (NH), where anomalies extend down into the lower troposphere and are concentrated over the Atlantic, resembling the negative phase of the North Atlantic Oscillation (NAO).

A more recent study by Zhang et al. (Submitted), that considered two models that differed only in their representation of interactive chemistry, also showed that changes in composition can impact the sign of the NH midlatitude jet response to increased CO_2. However, in contrast to CP2019, the long-term impact of this compositional feedback was a *poleward*, not equatorward, shift in the North Atlantic jet. Though not investigated in detail, this poleward shift of the jet was linked to changes in the ocean circulation, which were not examined in CP2019. More precisely, Zhang et al. (Submitted) noted that the AMOC exhibited a stronger decline in interactive simulations in which trace gases and aerosols were allowed to respond to increased CO_2, relative to non-interactive versions. Indeed, recent studies have highlighted the large influence that changes in the AMOC exert on the response of the NH midlatitude jet to increased CO_2 (Gervais et al. (2019)), with models featuring a larger AMOC decline also tending to produce a stronger poleward jet shift (Bellomo et al. (2021); Liu et al. (2020); Orbe et al. (Under Review)).

The results from Zhang et al. (Submitted) suggest that composition feedbacks on the NH midlatitude jet may depend on the response of the ocean circulation. However, that study did not examine the mechanism underlying the stronger AMOC response in the interactive chemistry simulations nor did it isolate the role of ozone from influences due to other trace gases and aerosols. To this end, here we hypothesize that the ozone-induced negative NAO wind anomalies reported in CP2019 provide a potential pathway through which stratospheric ozone changes can influence the AMOC. Variations in the jet – namely those resembling the NAO – have long been shown to influence variability of the AMOC through changes in wind stress (Marshall et al. (2001); Zhai and Marshall (2014)). Modified air-sea fluxes of heat, water and momentum associated with variations in the NAO alter vertical and horizontal density gradients in the subpolar gyre, inducing changes in deep water formation and the AMOC (e.g., Visbeck et al. (1998); Delworth and Dixon (2000)). This pathway via the NAO has been used to demonstrate how sudden stratospheric warmings influence the variability of heat flux anomalies into the ocean and ocean mixed layer depths in the North Atlantic (O'Callaghan and Mitchell (2014)) as well as the strength of the AMOC itself (Reichler et al. (2012)).

We begin by showing results from non-interactive and fully interactive chemistry global warming experiments produced with the new high-top coupled atmosphere ocean version of the NASA Goddard Institute for Space Studies (GISS) climate model that were submitted to the Coupled

Model Intercomparison Project Phase 6 (CMIP6) (Eyring et al. (2016)). We then show that the AMOC response in the interactive simulations is largely associated with changes in stratospheric ozone, not aerosols, using new experiments in which the stratospheric ozone response to 4xCO_2 is isolated from changes in other trace gases and aerosols. In particular, we show that our model captures the ozone-induced negative NAO-like pattern first reported in CP2019; in addition, we also find that ozone-driven changes in surface friction speed further weakens the AMOC, resulting in a long-term poleward shift of the NH jet. As a result, we show that both stratospheric ozone changes and the AMOC influence the NH jet on distinct "fast" and "long" timescales (and in the opposite sense), comprising a coupled atmosphere-ocean feedback on the NH midlatitude jet response to increased CO_2. While the former "fast" feedback was documented in CP2019, the latter has, to the best of knowledge, not been reported in previous studies.

It is important to note that previous studies have long shown that interactive atmospheric composition can strongly influence the AMOC, placing an almost exclusive focus on the role of aerosols Booth et al. (2012); Cowan and Cai (2013); Swingedouw et al. (2015). More recently, Rind et al. (2018) also identified a larger sensitivity of the AMOC response to global warming using an interactive configuration of the CMIP5 version of the GISS climate model (GISS-E2-R), compared to a non-interactive version. In that study, multicentennial cessations of the AMOC were found to occur in simulations in which natural aerosols (primarily sea salt) were allowed to locally cool sea surface temperatures through their influence on cloud optical thickness; these cooler SSTs were then linked to reduced evaporation relative to precipitation, resulting in positive surface freshwater forcing and reduced NADW production. As in Rind et al. (2018) we also show that compositional feedbacks play an important role on the response of the AMOC to CO_2 through their influence on surface fluxes and surface temperatures. However, the mechanism proposed here only invokes changes in stratospheric ozone, not aerosols.

We begin by discussing methods in Section 2 and present key results and conclusions in Sections 3 and 4, respectively.

2. Methods

a. Model and Configurations

Here we use the NASA Goddard Institute for Space Studies (GISS) "Middle Atmosphere (MA)" Model E2.2 (Rind et al. (2020); Orbe et al. (2020)). E2.2 consists of 102 vertical levels spanning the surface up to 0.002 hPa and is run at a horizontal resolution of 2 degrees by 2.5 degrees. Orographic and non-orographic gravity wave drag is parameterized following Lindzen (1987) and Rind et al. (1988), producing in E2.2 a quasibiennial oscillation (QBO) that compares well with observations as well as improved stratospheric polar vortex variability (Ayarzagüena et al. (2020); Rind et al. (2020)). Among the different model versions discussed in Rind et al. (2020) here we focus on the "Altered-Physics" (-AP) Version (E2.2-AP) because this is the configuration that was submitted CMIP6 and presented in recent studies (Ayarzagüena et al. (2020); DallaSanta et al. (2021a,b)).

We begin by showing the results reported in Zhang et al. (Submitted) using both "Non-INTeractive" (NINT) (Table 1, row 1-3) and fully interactive OMA ("One-Moment Aerosols"; Bauer et al. (2020)) configurations (Table 1, row 4-6). In the NINT configuration (denoted in CMIP6 as "physics version 1" on the Earth System Grid Federation (ESFG; https://esgf.llnl.gov)) all trace gases and aerosols are set to preindustrial values. Hence, in the 2- and 4xCO_2 NINT runs neither ozone nor other trace gases (besides water vapor) change in response to increased CO_2. By comparison, the OMA 2- and 4xCO_2 runs (denoted in CMIP6 as "physics version 3" on ESGF) capture the full nonlinear ozone response to CO_2, as well as composition feedbacks associated with other trace gases and aerosols.

In order to isolate the role of ozone feedbacks on the circulations, we then use a linearized ozone (LINOZ) configuration (Table 1, row 7-8). In LINOZ (McLinden et al. (2000)) the ozone field is calculated interactively by Taylor expanding the equation of state around present-day (2000–2010) values such that the ozone tendency is, to first-order, parameterized as a function of the local ozone mixing ratio, temperature, and overhead column ozone. Tropospheric ozone is calculated using monthly mean ozone production and loss rates archived from GEOS-CHEM (Rind et al. (2014)). In contrast to NINT, therefore, the LINOZ ensemble captures the influence of the ozone response to CO_2 on the large-scale circulation. Unlike OMA, however, it is much more computationally

TABLE 1. The Model E2.2 experiments presented in this study, including preindustrial control, abrupt $2xCO_2$ and abrupt $4xCO_2$ simulations using both NINT (rows 1-3) and OMA (rows 4-6) configurations. Four NINT abrupt $4xCO_2$ ensemble members are included (row 3) in order to compare with a four member $4xCO_2$ ensemble produced using the LINOZ configuration (row 8). The $4xCO_2$ ensemble mean LINOZ ozone response is also used to force four AMIP preindustrial experiments (row 9) in which all forcings other than ozone are set to preindustrial values. A LINOZ preindustrial control simulation (row 7) is also examined. All coupled simulations are run using the the GISS Ocean v1 (GO1) (i.e., "-G" in CMIP6 notation).

Configuration	Ozone	CO_2	Ensemble Size	SSTs and SICs
NINT	Preindustrial	Preindustrial	1	coupled (-G ocean)
NINT	Preindustrial	$2xCO_2$	1	coupled (-G ocean)
NINT	Preindustrial	$4xCO_2$	4	coupled (-G ocean)
OMA	Preindustrial	Preindustrial	1	coupled (-G ocean)
OMA	$2xCO_2$	$2xCO_2$	1	coupled (-G ocean)
OMA	$4xCO_2$	$4xCO_2$	1	coupled (-G ocean)
LINOZ	Preindustrial	Preindustrial	1	coupled (-G ocean)
LINOZ	$4xCO_2$	$4xCO_2$	4	coupled (-G ocean)
NINT	LINOZ $4xCO_2$	Preindustrial	4	AMIP (Preindustrial SSTs and SICs)

efficient to run and isolates the ozone feedback from feedbacks related to other trace gases and aerosols. DallaSanta et al. (2021a) previously showed that the LINOZ ozone parameterization reproduces well the vertical structure and seasonal cycle of stratospheric ozone obtained from the fully interactive OMA configuration (see their Figure 1).

b. Experiments

For the different model configurations (NINT, OMA, LINOZ) we perform 150-year-long abrupt 2- and $4xCO_2$ experiments, in which CO_2 values are abruptly doubled and quadrupled relative to preindustrial values. For each model configuration, these experiments are branched from a corresponding preindustrial control simulation. For NINT and LINOZ four-member $4xCO_2$ ensembles are run in order to assess the robustness of any ozone feedbacks. These experiments are all conducted using the atmosphere-ocean version of E2.2-AP that is coupled to the GISS Ocean v1 (GO1) (i.e., "-G" in CMIP6 notation, hereafter simply E2.2-G). For coupled atmosphere-ocean configurations in which (four-member) ensembles are run, different ensemble members are chosen from different initial ocean states spaced 20 years apart in the corresponding preindustrial control simulation.

In addition to the coupled atmosphere-ocean experiments, we also present results from a four-member ensemble of 60-year-long atmosphere-only AMIP experiments in which sea surface temperatures (SSTs) and sea ice concentrations (SICs) are fixed to preindustrial values, but the monthly mean time-evolving ensemble mean ozone response from the coupled LINOZ 4xCO_2 experiments is prescribed (Table 1, row 9). This allows us to quantify the impact of the ozone feedback represented in LINOZ on the large-scale circulation, absent any contributions from changes in background CO_2, sea ice concentrations or sea surface temperatures.

c. Analysis

1) Timescales

When examining the midlatitude jet response to increased CO_2 we account for the fact that extratropical circulation changes consist of distinct "fast" and "slow" responses (Ceppi et al. (2018), hereafter CZS2018). More precisely, CZS2018 show that most of the shift of the midlatitude jets occurs within 5-10 years of a steplike (abrupt) CO_2 forcing, with little shifts occurring during a slower response over which SSTs change over subsequent decades. In contrast to the Southern Hemisphere, zonal asymmetries play an important role in the Northern Hemisphere, where the influence of local patterns in sea surface temperature change can result in oppositely signed jet shifts on "slow" timescales. Given this potential for compensating jet shifts on distinct timescales, we therefore decompose the CO_2 circulation response into "fast" and "long" timescale responses.

More precisely, to account for the large internal variability in our runs, perhaps related to a somewhat larger ENSO amplitude in our model compared to observations (Rind et al. (2020)), we modify the original approach used in CZS2018 to define our "fast" response as the difference between the ensemble mean 4xCO_2 response, averaged over years 5-20 (as opposed to years 5-10), and the corresponding preindustrial control simulation. In addition, instead of focusing on the "slow" response, defined in CZS2018 as the difference between averages over years 120-140 and years 5-10, here we examine the "long" response, defined as the difference between the ensemble mean 4xCO_2 response, averaged over years 100-150, and the preindustrial control simulation. This approach is more consistent with what was used in Zhang et al. (Submitted) and CP2019, with which we directly compare our results throughout. Note that in response to an abrupt quadrupling of CO_2 the NINT model configuration produces global mean surface temperature "fast" and "long" responses of ~2.9°C and ~3.9°C, respectively. Statistical significance of all changes are assessed relative to the interannual variability in the corresponding preindustrial control simulation for each configuration (Table 1, rows 1,4,7).

2) Analysis Fields

In addition to the atmospheric variables examined in CP2019 (i.e., zonal mean wind, zonal mean temperature, surface temperature, 850 hPa zonal wind) we examine ocean variables relevant to understanding the evolution of the AMOC and its coupling to the atmosphere. In particular, in addition to examining the surface mixed layer depths we also examine sea surface temperatures, surface friction speed, horizontal ocean heat and salinity transports as well as the net heat fluxes which, together with the net freshwater fluxes, F (inferred from precipitation minus evaporation (P-E)), provide information about the surface buoyancy forcing (Large and Yeager (2009)). In our simulations, the preindustrial climatological buoyancy forcing over the North Atlantic is dominated by the net heat fluxes ($Q = Q_H+Q_E+Q_S+Q_L$), which are defined to be positive into the ocean (Appendix Figure 1, left). These are further partitioned into their respective latent heat (Q_E) and sensible heat (Q_H) contributions as we find that the net solar (Q_S) and longwave (Q_L) flux radiative contributions are negligible over the North Atlantic region (Appendix Figure 1, right).

Given our interest in the Northern Hemisphere we focus primarily on December-January-February (DJF). The ocean heat transport changes in our simulations are also most pronounced during DJF, consistent with the analyses presented in Romanou et al. (Under Review) and Orbe et al. (Under Review).

3. Results

a. Abrupt 2xCO_2 and 4xCO_2 Zonal Mean Wind Response: OMA versus NINT

Before focusing on ozone feedbacks, we first review the OMA versus NINT differences in NH jet behavior that were presented in Zhang et al. (Submitted) (Figure 1). In the stratosphere the zonally averaged DJF wind response to 2- and 4xCO_2 features an acceleration at nearly all latitudes, consistent with amplified warming in the tropical upper troposphere (Shaw (2019)) and increased

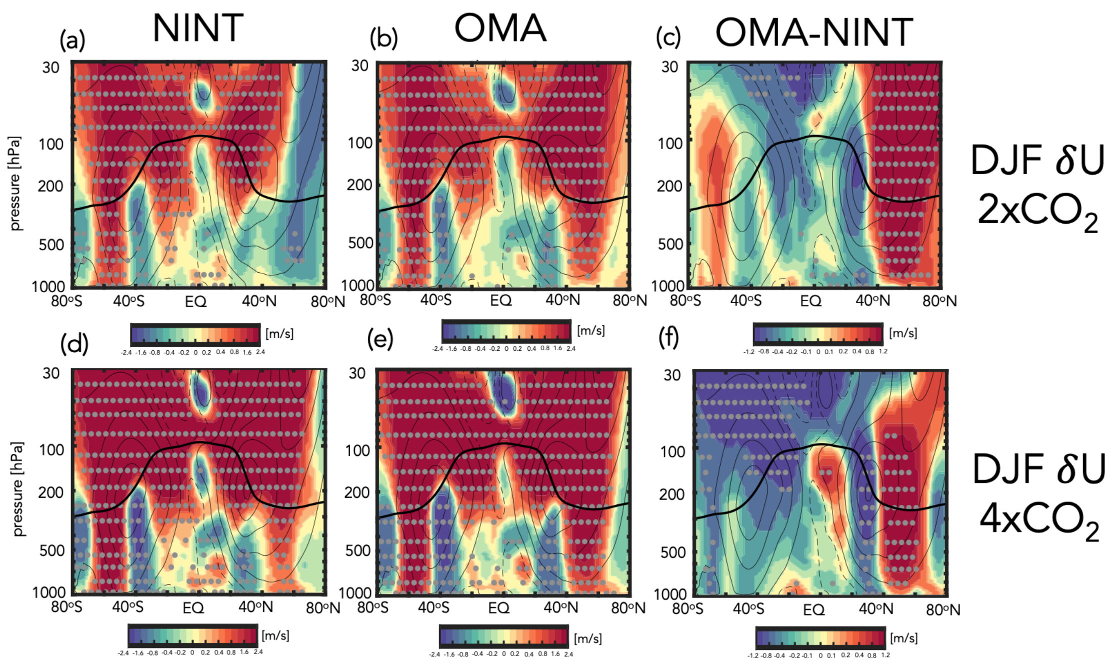

FIG. 1. Colors show the December-January-February (DJF) response of the zonal mean zonal winds, U, to an abrupt doubling (top) and quadrupling (bottom) of CO_2, averaged over years 100-150. Results are shown for the "Non-INTeractive" (NINT) (a,d) and fully interactive OMA ("One-Moment Aerosols") configurations (b,e), where one ensemble member has been used for each forcing scenario. The OMA - NINT differences are also shown (c,f). Black contours denote climatological mean DJF U values (contour interval: 8 m/s). Stippled regions are statistically significant and the black thick line shows the climatological mean tropopause in the preindustrial control NINT simulation. Note that all colorbar bounds are consistent with those use in Chiodo and Polvani (2019) in order to facilitate comparisons with that study.

cooling of the stratosphere with height (Garcia and Randel (2008)). Similar wind responses emerge in both the NINT and OMA configurations, except over northern high latitudes at $2xCO_2$, where the strengthened zonal winds in NINT are not statistically significant.

In the troposphere, however, there are noticeable differences between the OMA and NINT simulations. In particular, the NH midlatitude jet features a much stronger poleward shift in OMA, compared to NINT (Figures 3 and 6 in Zhang et al. (Submitted)). As discussed in that study, the stronger response in OMA results in enhanced eddy mixing along isentropes on the poleward flank of the NH jet, resulting in increased transport of tracers from the northern midlatitude surface to the Arctic (not shown). This difference between OMA and NINT occurs at both 2- and at $4xCO_2$,

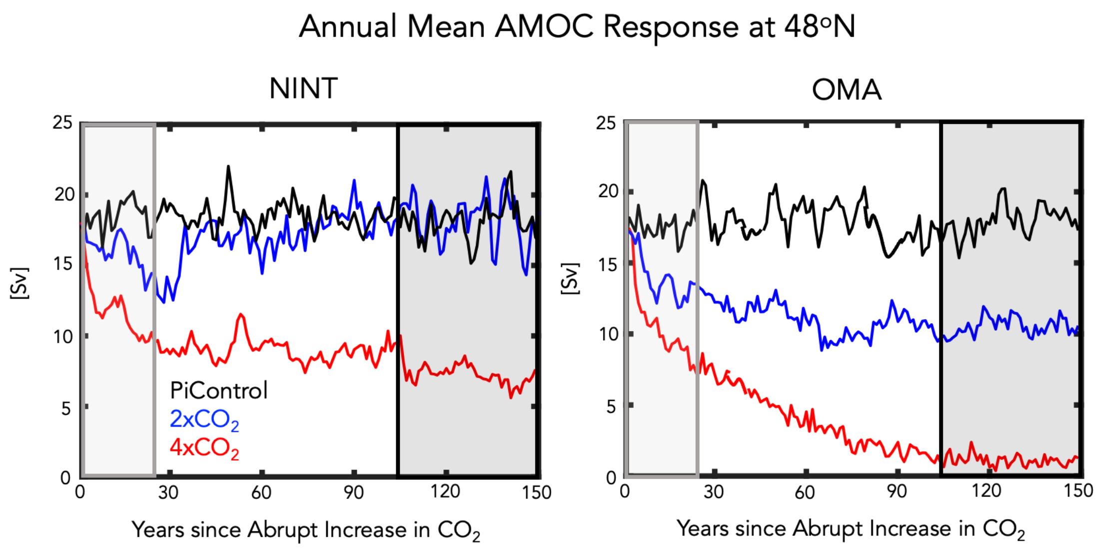

FIG. 2. Changes in the annual mean maximum overturning stream function in the Atlantic ocean, evaluated at 48°N, for the preindustrial control (black), abrupt 2xCO_2 (blue) and abrupt 4xCO_2 (red) simulations. Results for the NINT (left) and OMA (right) configurations are shown. Light grey and black shaded boxes denote the "fast" and "long" timescale response averaging periods.

resulting in a nonlinearity in the jet (and tracer transport) response in NINT that is not present in the OMA simulations. In the SH, by comparison, the differences between OMA and NINT are much smaller and not statistically significant.

Zhang et al. (Submitted) showed that the nonlinearity in NH jet behavior evident in the "long" response in the NINT model configuration was related to a nonlinear AMOC response to CO_2 forcing (Figure 2). That is, despite an initial weakening, in response to 2xCO_2, the AMOC eventually recovers in the NINT 2xCO_2 simulation to preindustrial values, in contrast to the response to 4xCO_2 in which the AMOC is about 10 SV weaker than the preindustrial control (black boxes). This results in a so-called "AMOC nonlinearity" of ~-5SV in the NINT configuration. By comparison, in the OMA configuration, the AMOC weakens by ~7 and ~17 SV in the 2- and 4xCO_2 simulations, respectively, representing only a very weak nonlinearity in the AMOC (of ~1.5 SV).

As it is difficult to meaningfully interpret the zonal mean wind response in the NH, where there are large zonal variations in the midlatitude jet (Simpson et al. (2014)), we next compare the 850 hPa zonal wind changes between the NINT and OMA 4xCO_2 simulations, further distinguishing

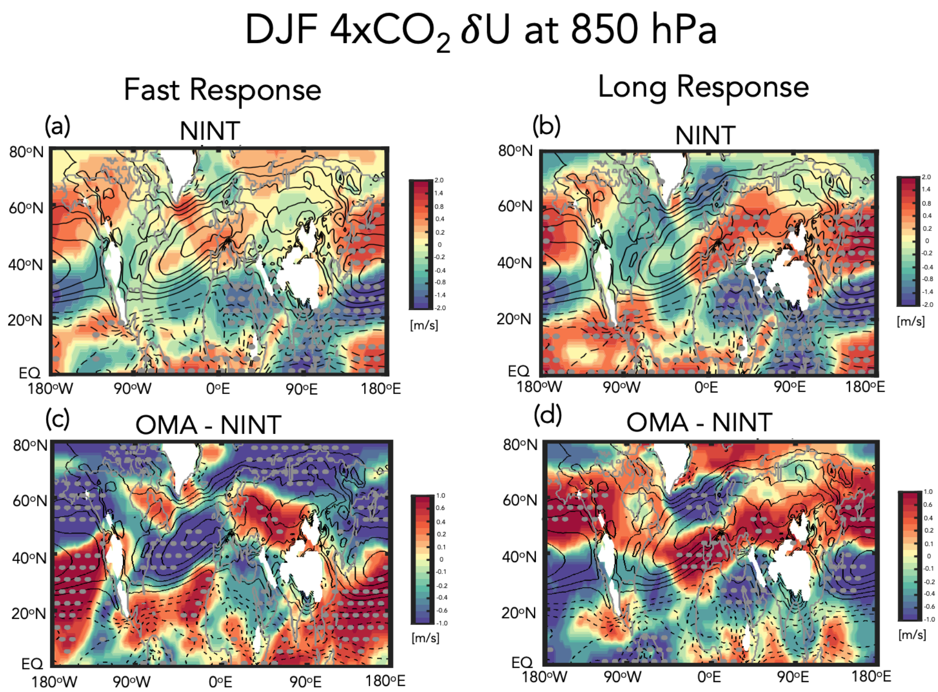

Fig. 3. Colors show the $4xCO_2$ (four member) ensemble mean change in the DJF 850 hPa zonal winds for the NINT configuration, decomposed into "fast" (i.e. years 5-20) (a) and "long" (i.e. years 100-150) (b) responses. The OMA - NINT fast and long differences are shown in (c) and (d), respectively. Note that one ensemble member is used in displaying the OMA - NINT differences (same as used in Figure 1). Black contours denote climatological mean DJF values (U contour interval: 2 m/s) and stippled regions are statistically significant.

between "fast" and "long" responses (Figure 3). We begin with the NINT equilibrated or "long" response (i.e. years 100-150), which consists of a poleward jet shift over the Pacific basin and an acceleration and eastward extension of the jet over the Atlantic (Fig. 3b). This pattern is amplified in the OMA run (Fig. 3d), in which both the strengthening of the jet over the Atlantic and its poleward shift over the Pacific are more pronounced. This wind response in OMA, relative to NINT, is consistent with the jet differences identified in Orbe et al. (Under Review) between two non-interactive simulations of the GISS low-top climate model in which only the AMOC strength differed. This suggests that the jet differences between OMA and NINT on these longer timescales are primarily driven by differences in the AMOC response, as concluded in Zhang et al. (Submitted).

Figure 2 (grey boxes) highlights how the AMOC differences between OMA and NINT noted in Zhang et al. (Submitted) arise very early in the simulations (within the first 20 years). Over these years – which comprise the "fast" response – the impact of interactive chemistry on the zonal wind changes is very different (Fig. 3a,c). In particular, over the Atlantic, interactive composition results in a strong weakening over the jet core and an acceleration on the equatorward flank of the jet (Fig. 3c). The jet response is also very different over the Pacific, where the jet shifts equatorward, not poleward as in the NINT simulation (Fig. 3a).

This fast composition feedback that occurs over years 5-20 is consistent with the results from CP2019, who showed that the ozone response to 4xCO_2 induces a weakening of the North Atlantic jet and a strengthening on its equatorward flank (see their Figure 6). This response is reminiscent of the negative phase of the NAO which previous studies have shown can result in a weaker AMOC (Delworth and Zeng (2016)). In CP2019, however, this response is realized through changes in stratospheric ozone alone, whereas in OMA all trace gases and aerosols are responding. Furthermore, the significance of this rapid response with only one ensemble member is uncertain, particularly during the first 5-20 years when the signal is confounded by large internal variability. To this end, next we present results from the larger (4-member) LINOZ ensemble to examine whether the fast response in the NH jet is related to stratospheric ozone changes.

b. Abrupt 4xCO_2 Stratospheric Ozone and Temperature Responses: OMA versus LINOZ

Before examining the circulation response in the LINOZ ensemble, we first compare the annually averaged ensemble mean LINOZ 4xCO_2 ozone response with that from the OMA simulation (Figure 4). The amplitude and pattern of the ozone response in the LINOZ ensemble (Fig. 4b) is generally very similar to the ozone response in the OMA simulation (Fig. 4a). In both configurations the pattern of the 4xCO_2 changes reflects a decrease in tropical LS ozone, associated with enhanced tropical upwelling (Garcia and Randel (2008)), and enhanced concentrations over high latitudes. Over all latitudes the ozone changes are statistically significant, relative to interannual variability in the preindustrial control simulation.

Over northern high latitudes there are some differences in the mid-to-lower stratosphere (~30-100 hPa) between LINOZ and OMA, generally consistent with Chiodo et al. (2018), who found that in this region the ozone response to CO_2 is somewhat more model dependent. Furthermore, both simulations feature small changes in the troposphere. Overall, therefore, the LINOZ scheme captures the gross characteristics of the ozone abrupt 4xCO_2 response expected from previous

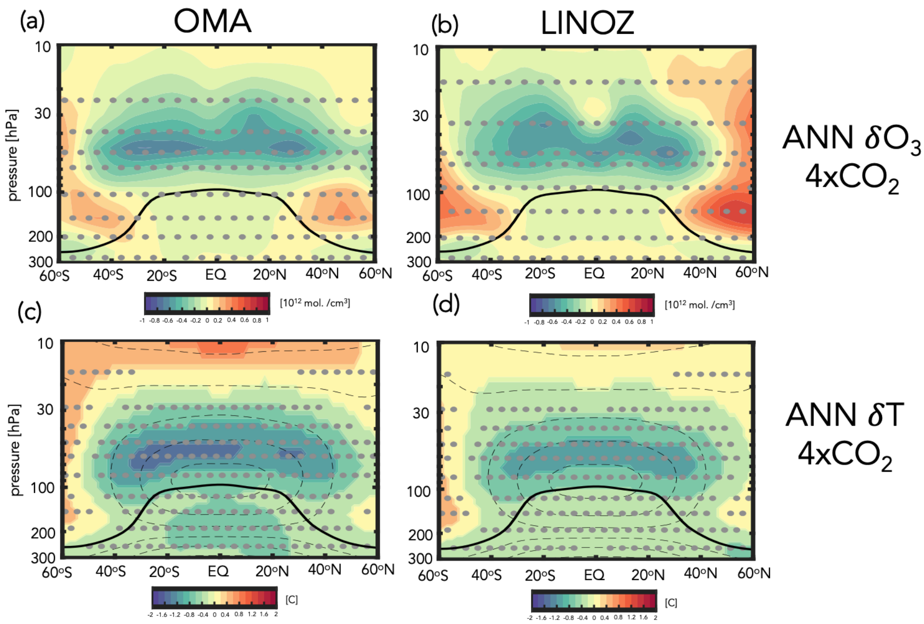

FIG. 4. Colors show the annual averaged change in ozone number density (top) and temperature (bottom) in response to $4xCO_2$. Results for OMA (left) and LINOZ (right) are shown, averaged over years 5-20. One simulation is shown for OMA and the four-member ensemble mean response is shown for LINOZ. Black contours in the bottom panels show climatological mean temperatures (contour interval: 10 C). Stippled regions are statistically significant and the black thick line shows the climatological mean tropopause in the preindustrial control NINT simulation.

studies. Note that this ozone response occurs in both simulations within the 5-20 years that comprise the "fast" response timescale, although full equilibration at high latitudes does take somewhat longer (not shown).

In response to the ozone changes to $4xCO_2$ both the OMA simulation and LINOZ ensemble produce cooling in the tropical lower stratosphere and warming over high latitudes (Fig. 4c,d). The amplitude of the cooling is ~3K in the tropical lower stratosphere, and is more-or-less collocated with the region of largest ozone decreases. Further analysis of the temperature tendencies reveals that in our model the cooler temperatures in the tropics (20°S-20°N) and high latitudes (> 40°N) are respectively associated with reduced and increased radiative heating, primarily in the shortwave component (not shown). Dynamically, comparisons of the $4xCO_2$ changes in the residual mean stream function show a weaker response in LINOZ, relative to NINT (not shown). This ozone feedback on the Brewer-Dobson circulation, first identified in DallaSanta et al. (2021a), contributes to reduced upwelling, adiabatic cooling, and ozone transport within the lower tropical stratosphere. These circulation changes are therefore not the primary drivers of the temperature response; rather, they are primarily determined by the shortwave radiative response to ozone changes (CP2019).

Despite the somewhat stronger cooling in OMA (Fig. 4c) compared to NINT (Fig. 4d), the temperature response in both configurations is within the 2-4 K range documented in CP2019 (note that all colorbars used are consistent with that study to facilitate comparisons with their results). An important point to note is that the temperature changes due to ozone are of a similar magnitude to the temperature changes due to 4xCO_2 alone in the tropical lower stratosphere (i.e., considering no ozone feedback), where the stratosphere cools by ~2K in the NINT ensemble (not shown). The ozone changes present in LINOZ (and OMA) therefore represent a substantial feedback on the CO_2-induced cooling in the stratosphere.

c. Ozone Feedback on Northern Hemisphere Jet: Fast Response

The temperature response due to ozone is dynamically consequential for the troposphere to the extent that it modifies temperature gradients (and winds) in the lower stratosphere. Indeed, the LINOZ ensemble shows a strong enhancement of lower stratospheric temperature gradients in both hemispheres on both the fast and long response timescales (Fig. 5a,b). In the fast response, this reduction in the meridional temperature gradient near the tropopause has important consequences for the midlatitude jet in both hemispheres, particularly in the NH where it strengthens above and along the jet core and weakens on the poleward flank of the jet over latitudes north of ~ 50°N (Fig. 5c). The winds also accelerate equatorward of the jet core, relative to NINT, in both hemispheres, although the response is only statistically significant in our model in the NH. This ozone-induced response in the jet is very similar to the pattern of the wind response reported in CP2019 (see their Figures 4 and 5). As with the temperature changes occurring in the lower stratosphere, the wind response to ozone changes is similar in magnitude to the 4xCO_2 response, again suggesting a substantial modulation of the circulation in both hemispheres by ozone changes alone.

The fast zonal mean response to ozone changes reflects a weakening of the polar jet over all longitudes, with the largest negative anomalies concentrated over the Atlantic ocean which are flanked equatorward by positive wind anomalies (Fig. 6a). These wind changes are vertically coherent throughout the troposphere as the LINOZ-NINT changes are similar at 300 hPa (not

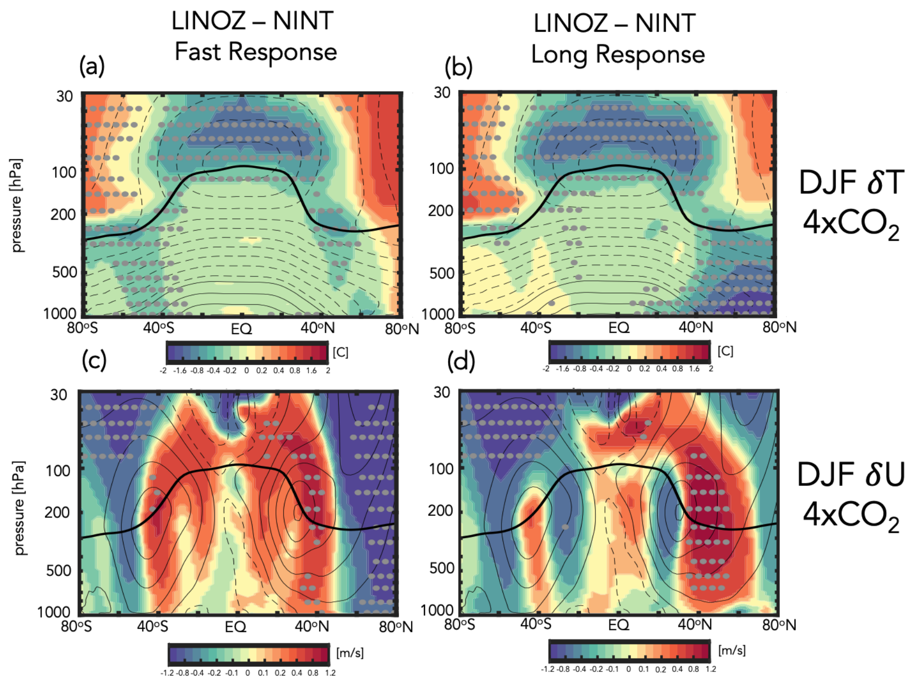

FIG. 5. Colors show the LINOZ-NINT ensemble mean difference in the DJF response of the zonal mean temperatures, T (top) and zonal winds, U (bottom) in response to an abrupt quadrupling of CO_2. Both LINOZ and NINT ensembles consist of four members. Responses are decomposed into "fast" (a,c) and "long" (b,d) changes. Contours denote climatological mean DJF values (T contour interval: 10 C; U contour interval: 8 m/s). Stippled regions are statistically significant and the black thick line shows the climatological mean tropopause in the preindustrial control simulation.

shown). This LINOZ-NINT wind dipole over the Atlantic is very similar to the fast wind response captured in the fully interactive OMA simulation (Fig. 3c), especially over the Atlantic. Over the Pacific, by comparison, the OMA and LINOZ responses are different, consistent with CP2019 who found no robust ozone feedback over the Pacific (see their Figure 5). Furthermore, the weakening of the North Atlantic jet in the LINOZ simulations is associated with warming over North America and cooling over the North Atlantic and over Eurasia, resembling the negative phase of the NAO (Fig. 6c). A similar surface temperature anomaly was identified in CP2019 (see their Figure 7) and in our model occur in conjunction with positive sea level pressure (SLP) anomalies over the Arctic (Appendix Figure 2, top), both features being reminiscent of a negative NAO.

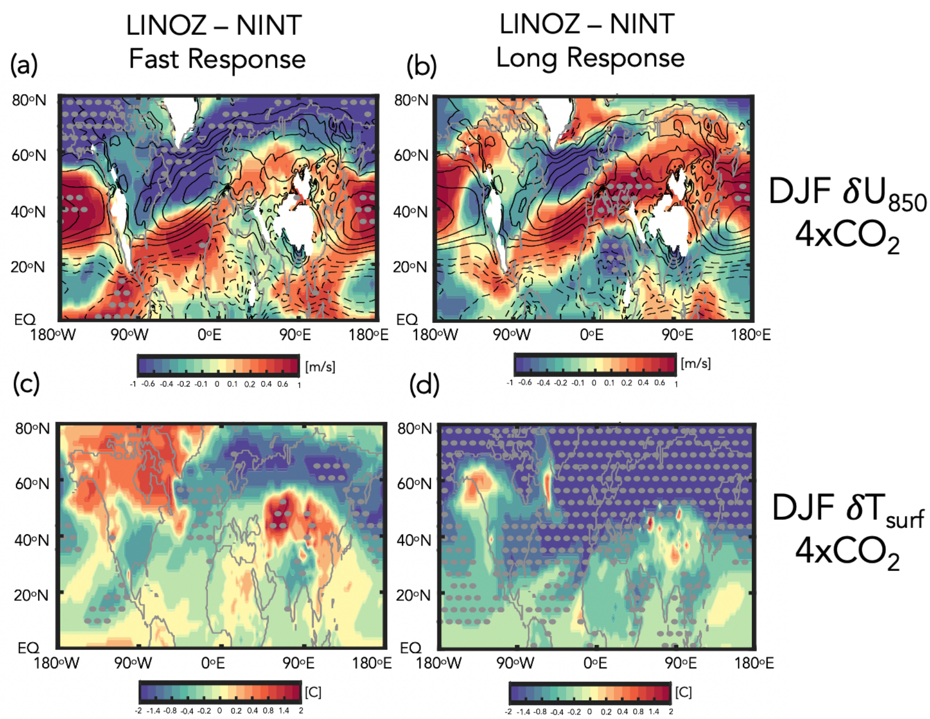

FIG. 6. Same as Figure 5, except showing the LINOZ-NINT DJF response in the 850 hPa zonal winds, U_{850} (top) and surface temperatures, T_{surf} (bottom). Contours in top panels denote climatological mean DJF values of U_{850} (contour interval: 2 m/s). Note the similarity between the "fast" wind response shown in (a) and the CP2019 results (their Figure 6).

d. Ozone Feedback on Northern Hemisphere Jet: Long Response

Interestingly, while the fast responses in the winds and temperatures in the LINOZ ensemble are highly consistent with the results from CP2019, our model also simulates a distinct "long" response characterized by strong cooling over the Arctic from the surface to the mid-to-upper troposphere (Fig. 5b). This cooling, which was not identified in CP2019, results in enhanced mid-to-lower tropospheric temperature gradients, prompting a strong poleward shift of the NH jet and a statistically significant acceleration of the winds at 50°N exceeding 2 m/s (Fig. 5d).

Zonally, the cooling over the Arctic occurring in the LINOZ ensemble during the long response primarily reflects hemispheric-wide cooling over the Arctic associated with an expansion of the North Atlantic Warming Hole (Fig. 6d). This enhancement of meridional temperature gradients in the lower and mid troposphere drives a poleward shift that spans all longitudes and originates over the North Atlantic (Fig. 6b), where the jet exhibits a distinct acceleration and eastward extension over Europe. Note that over the jet core (40°N-50°N) the winds accelerate (in the zonal mean) during both "fast" (Fig. 5c) and "long" responses (Fig. 5d). However, north of 50°N the responses are very different, with the fast response exhibiting a strong weakening, in contrast to the acceleration ocurring on longer (i.e., "long" response) timescales. This behavior north of 50°N was not captured in CP2019 and comprises a coupled ozone-ocean feedback that is distinct from what was outlined in that study.

e. Long Ozone Feedback: Modulation by the AMOC

The "long" responses in the tropospheric winds and temperatures that occurs in the LINOZ ensemble are not obviously linked to ozone-driven temperature changes in the stratosphere, which do not extend into the troposphere. What, then, is the driver of the lower tropospheric high latitude cooling, if it is not directly linked to ozone-driven stratospheric temperature changes?

As expected from the OMA and NINT results presented in Zhang et al. (Submitted), we find that the strong cooling that occurs over the NH in the long LINOZ response is also related to a weakening of the AMOC at 4xCO_2 (Mitevski et al. (2021); Rind et al. (2020); Orbe et al. (Under Review)). In particular, Figure 7 shows stronger weakening of the AMOC in the LINOZ (green lines) ensemble, relative to NINT (blue lines) at both 26°N (left) and at 48°N (right). Despite large internal variability, the LINOZ ensemble shows a more rapid decline of the AMOC, a difference that is evident at both latitudes.

Interestingly, comparisons of the AMOC behavior in LINOZ with the fully interactive OMA simulation (red line) shows a striking similarity (and the mechanism of these changes is also similar, as shown in Section 3f). This similarity is surprising, given that other (non-ozone) trace gases and aerosols are also evolving in the OMA experiment. In particular, Rind et al. (2018), using a previous version of the model, observed an indirect effect of natural aerosols (primarily sea salt) on AMOC stability. They showed that aerosols enhanced the local cooling of SSTs in regions of increased cloud cover in a warmer climate by acting as condensation nuclei and thereby raising cloud optical thickness and ocean surface cooling. This surface cooling was then linked to reduced evaporation relative to precipitation, resulting in anomalously positive surface freshwater forcing and reduced North Atlantic Deep Water (NADW) production. That study, however, focused on aerosol-induced AMOC cessations occurring on multicentennial timescales

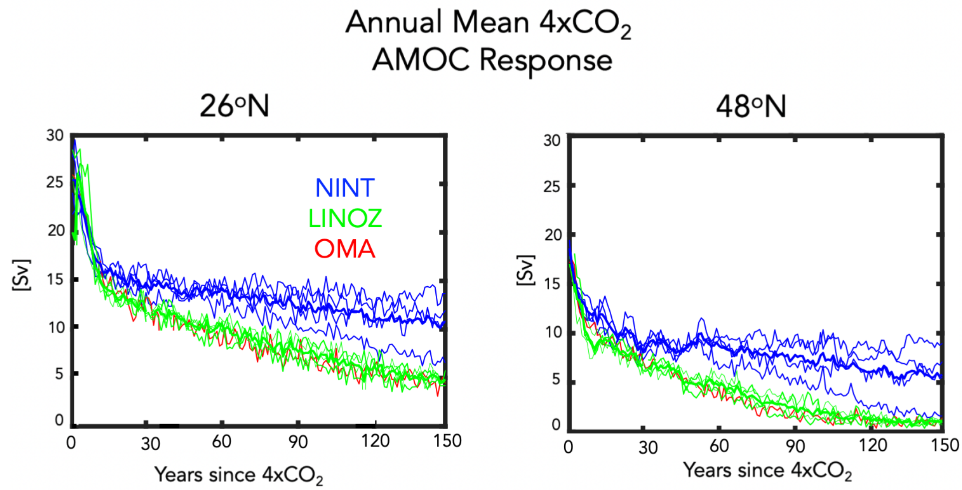

FIG. 7. Changes in the annual mean maximum overturning stream function in the Atlantic ocean, evaluated at 26°N (left) and 48°N (right) in response to 4xCO_2, relative to the preindustrial control simulations. Results for the LINOZ and NINT ensembles are shown in green and blue, respectively (thick lines denote ensemble means). Red lines show the response in the OMA simulation.

long after the initial (abrupt) warming. By comparison, the results in Figure 7 identify an impact of ozone on the AMOC that occurs within the first 20 years of the initial CO_2 forcing – that is, over the period during which stratospheric temperature gradients are most impacted by ozone (not aerosols). Our results, therefore, highlight that during this time frame the AMOC can be as (if not more) sensitive to wind-driven buoyancy changes forced by stratospheric ozone anomalies as they are to aerosol-induced changes in freshwater forcing.

Before elucidating the mechanism of the AMOC changes in the LINOZ ensemble, we first identify the region over which the largest differences in mixed layer depth begin to emerge between the LINOZ (OMA) and NINT simulations. In particular, the weaker AMOC in the LINOZ and OMA runs is found to be accompanied by a rapid reduction in mixed layer depths, which occur primarily in the Irminger Sea region (55°N-65°N, 40°W-20°W) (Figure 8). The mixed layer depth differences in the Labrador Sea are, by comparison, negligible. East of the Irminger Sea (i.e., 55°N−65°N, 20°W-0°) we also identify differences between the ensembles (not shown), but these emerge later, suggesting that the Irminger Sea changes are likely the initiators of the differences in AMOC behavior between the NINT and LINOZ ensembles. A similar region was identified in Romanou et al. (Under Review) as being key for determining the sensitivity of the AMOC in

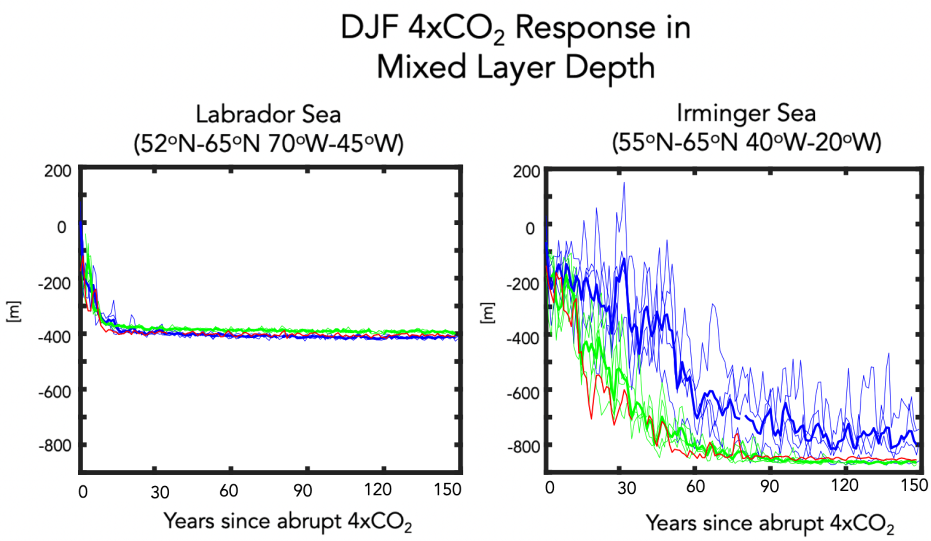

FIG. 8. Changes in the DJF mixed layer depths, evaluated over the Labrador Sea (left) and Irminger Sea (right) in response to $4xCO_2$, relative to the preindustrial control simulations. Results for the LINOZ and NINT ensembles are shown in green and blue, respectively (thick lines denote ensemble means). Red lines show the response in the OMA simulation.

various SSP 2-4.5 ensemble runs, albeit for simulations conducted using the low-top GISS climate model.

f. Ozone Feedback Dependence on the AMOC: Linking Fast and Long Responses

Is the fact that the AMOC declines more rapidly in the LINOZ ensemble – and the OMA run – a response to the ozone changes in those simulations or just a random occurrence? In the fast response the zonal wind changes over the North Atlantic reflect a weakening of the jet core that is flanked equatorward by positive anomalies, resembling a negative NAO pattern. Indeed, a negative (positive) NAO has been associated with a weaker (stronger) AMOC by adding (extracting) heat to/from the subpolar gyre, resulting in reduced (increased) NADW formation (Delworth and Zeng (2016)). Here we argue that such a mechanism is present in our model simulations, resulting in an additional substantial modulation of the NH midlatitude jet location by ozone, this time via its influence on the AMOC.

In particular, Figure 9 shows maps of the surface zonal wind, surface friction speed, mixed layer depth, net heat fluxes, sea surface temperatures, and north-south heat and salinity ocean transports over years 1-5. In response to an abrupt quadrupling of CO_2, there is a weak acceleration of the surface zonal winds on the poleward flank of the North Atlantic jet (~60°N-70°N) (Fig. 9a, top). Over the subpolar North Atlantic the surface winds weaken, leading to a significant reduction in surface friction speed (Fig. 9b, top) and mixed layer depths (Fig. 9c, top), as well as increased heat flux into the ocean (in the form of reduced latent heat fluxes out of the ocean) (Fig. 9d, top) and warmer sea surface temperatures (Fig. 9e, top). At these early years the changes in meridional heat and salinity transports over the Irminger Sea are relatively small (Fig. 9fg, top).

In response to the ozone changes captured in the LINOZ ensemble during years 1-5, there is a strong reduction in the surface zonal winds and friction speed (Fig. 9 ab, bottom), consistent with the negative NAO response evident in the 850 hPa zonal winds (Fig. 6c, top). The surface friction changes align closely with the reduced mixed layer depths which extend well into the Irminger Sea region and over latitudes further south of the subpolar gyre (Fig. 9c, bottom).

The reductions in mixed layer depth that occur over the Irminger Sea are likely driven by the reductions in surface wind speed which increased (primarily latent) heat fluxes into the ocean (Fig. 9d, bottom), driving warmer sea surface temperatures in LINOZ, relative to NINT (Fig. 9e, bottom). The sign of the response of the heat fluxes in the subpolar gyre region is consistent with previous studies showing that a positive (negative) phase of the NAO implies reduced (enhanced) atmosphere to ocean heat fluxes (Delworth et al. (2017)). Furthermore, the spatial pattern of the heat flux response is very similar to the NAO heat flux composites that were prescribed in Delworth and Zeng (2016) and inferred from observations in Ma et al. (2020) (see their Figure 6), who showed that there is much greater heat loss from the ocean over the subpolar region in association with a jet strengthening.

At the same time, the changes in freshwater forcing (P-E) during this time period are negligible such that the net buoyancy forcing (~Q+F) is positive. This stabilizing buoyancy forcing from surface warming makes the mixed layer depths shallower by suppressing convective mixing, shutting down NADW production (Alexander et al. (2000); Kantha and Clayson (2000)). There is also an initial change in the north-south heat and salt transports that is colocated with the dipole anomaly in the surface friction speed, promoting anomalous poleward salt and heat transport into the subpolar gyre (Fig. 9fg, bottom). This feature is confined to the top few ocean layers (not shown) and the implied anomalous heat transport could be contributing to the warmer sea surface temperatures in that region, in addition to the surface heat flux changes.

Over the ensuing years (5-20) a similar pattern is maintained in the LINOZ ensemble (Figure 10, middle row). The reduction in NADW, however, results in reduced northward heat and salinity transports (Fig. 10 fg, middle) throughout the ocean columm. While this results in cooler SSTs south of the subpolar gyre region (Fig. 10e, middle), which otherwise might enhance the density of the near-surface water masses, the reduced northward salinity transports prevent the AMOC from restarting. Interestingly, the results from the OMA simulation show a very similar response as the LINOZ ensemble (Figure 10, bottom row), suggesting that stratospheric ozone changes in that simulation are also likely the primary driver of the weaker AMOC in that model configuration. This sequence of processes linking the surface wind changes to anomalous heat fluxes and reduced NADW is basically identical to what is outlined in Figure 4 of Delworth and Zeng (2016) and Figure 1 of Khatri et al. (2022). Additional analysis of the 2xCO_2 simulations, which feature a stronger AMOC decline in OMA (and LINOZ) compared to NINT (Figure 2), reveals that a similar mechanism for reduced NADW production occurs at lower CO_2 forcing (not shown).

Finally, examining the timescale of the responses of the variables shown in Figures 9 and 10 reinforces the strong coupling between the changes in surface friction speed, sea surface temperature, latent heat fluxes and mixed layer depth changes over the Irminger Sea region (Figure 11a-d). Despite large internal variability, there is a clear separation between the LINOZ (and OMA) and NINT ensembles that emerges around year 15 (black dashed lines). The changes in sensible heat emerge after the latent heat fluxes (Fig. 11e), suggesting that the latter play a more important contribution in initializing the heat flux differences in LINOZ (and OMA), relative to NINT.

Finally, while they may contribute to enhanced positive buoyancy forcing later in the integrations, the freshwater forcing anomalies (F = P-E) are shown to be negligible during the initial years following the abrupt quadrupling of CO_2 (Fig. 11f), indicating that the primary driver of the initial difference between the LINOZ (and OMA) and NINT runs is related to the surface wind-driven changes as they impact the latent heat fluxes into the ocean. This is consistent with Roach et al. (2022) who showed a much stronger correlation between AMOC strength at 26°N and the heat component of the surface buoyancy flux, relative to the freshwater component, in various

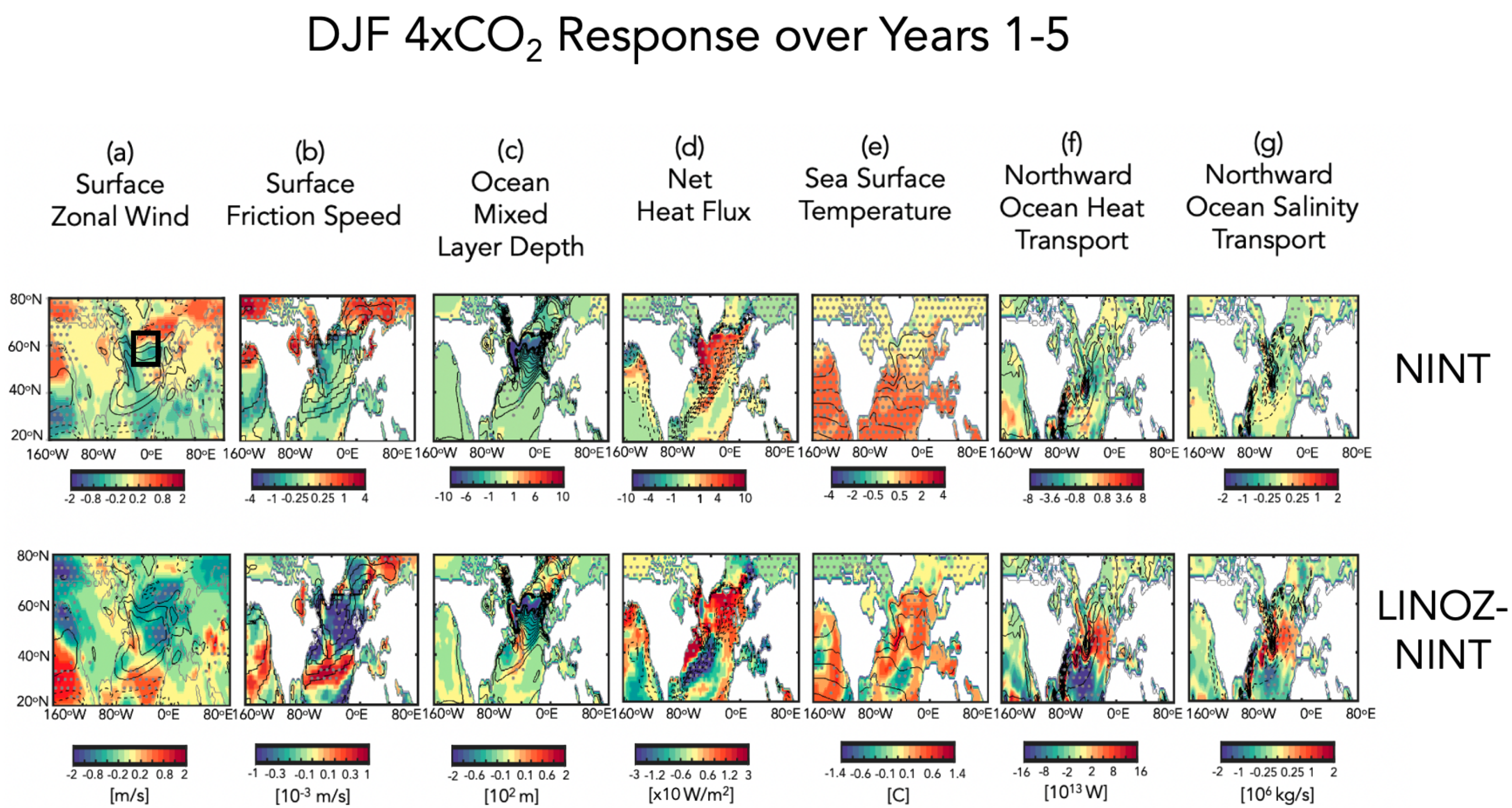

FIG. 9. Top panels: Colors show the December-January-February (DJF) response of the surface zonal wind (a), surface friction speed (b), ocean mixed layer depth (c), net heat flux (sum of sensible plus latent heat) (d), sea surface temperature (e) and northward heat (f) and salt (g) transports in response to an abrupt quadrupling of CO_2. Results are shown for the 4-member ensemble averaged NINT configuration. Bottom panels: Same as top panels, except showing the LINOZ minus NINT ensemble mean difference. For both top and bottom panels, responses have been averaged over years 1-5 since "branching" from the preindustrial control simulation. Stippled regions are statistically significant and black contours denote climatological mean DJF values. Contour intervals: surface zonal wind [2 m/s], surface friction speed [2.5×10^{-3} m/s], mixed layer depth [60 m], net heat flux [30 W/m^2], sea surface temperature interval [2 C], northward heat flux [2×10^{12} W], and northward salt flux [10^6 kg/s]. The black box in (a) bounds the Irminger Sea region over which the spatial averages in Figure 8b and Figure 11 are evaluated.

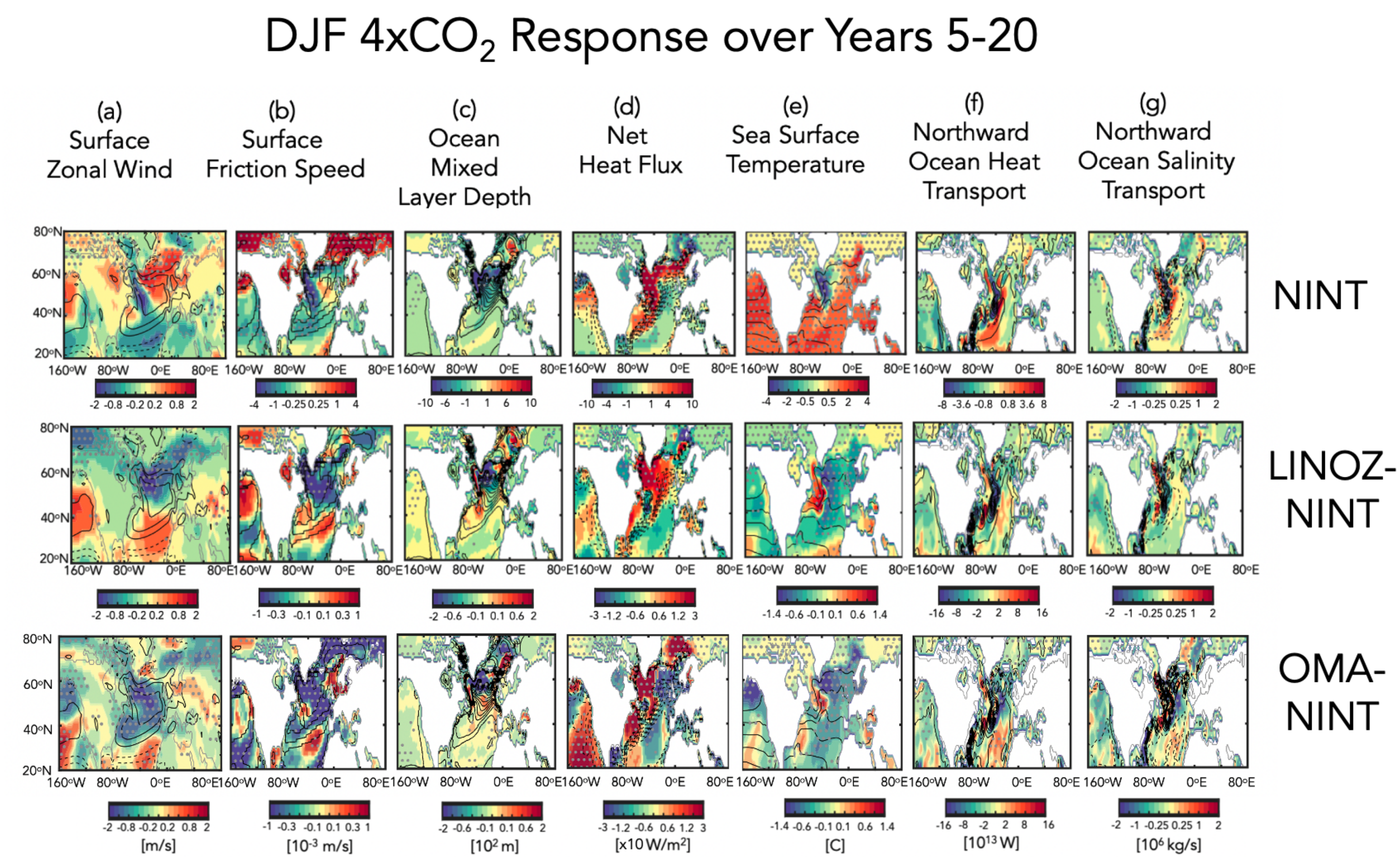

FIG. 10. Same as Figure 9, except showing the responses, averaged over years 5-20. An extra row at the bottom has been added, showing the OMA - NINT differences, where the ensemble members shown in Figures 1, 2 and 3 have been used. Same contour intervals and colorbars have been used as in Fig. 9.

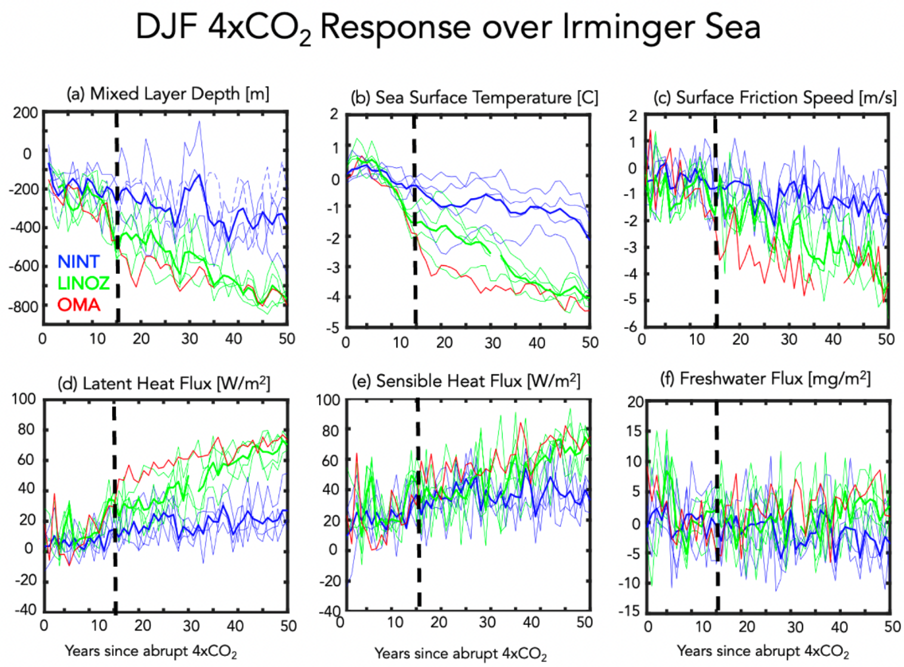

FIG. 11. Changes in the DJF mixed layer depths (a), sea surface temperatures (b), surface friction speed (c), latent heat fluxes (d), sensible heat fluxes (e) and precipitation minus evaporation (f) in response to 4xCO_2, relative to the preindustrial control simulations. Averages are over the Irminger Sea (55°N-65°N, 40°W-20°W). Results for the LINOZ and NINT ensembles are shown in green and blue, respectively (thick lines denote ensemble means). Red lines show the response in the OMA simulation. Black vertical lines indicate year ~15 at which point the mixed layer depth responses in the LINOZ and NINT ensembles diverge. Note that the freshwater flux unit of 1 mg/m^2 per second (≡ 0.0864 mm/day ≡ 3.1 cm/year) is used, because at 5°C it contributes approximately the same ocean density flux as the heat flux unit of 1 W/m^2 (Large and Yeager (2009)).

experiments using the Community Earth System Model version 1 (CESM1) in which the winds over the subpolar gyre were nudged to reanalysis values. Note that in our model other potential contributors to freshwater forcing from sea ice do reveal differences between the LINOZ, OMA and NINT ensembles, but these emerge several years (i.e., years ~20-30) after the changes in sea surface temperatures and heat fluxes (not shown).

g. *Ozone Driver of AMOC Changes: Fixed SST Results*

So far, we have shown that the stratospheric ozone changes that occur in response to 4xCO_2 result in a negative NAO response over the North Atlantic (Fig. 5,6). In our model this triggers a more rapid decline of the AMOC (Fig. 7) through surface-wind driven changes in heat fluxes into the ocean (Fig. 9,10). While the time series analysis (Fig. 11) reveals that the AMOC changes in the LINOZ (OMA) ensemble occur on similar timescales as the wind (and heat flux) changes, one potentially confounding factor is the fact that the AMOC reduction itself results in reduced wind speeds over the subpolar gyre region. These reduced near-surface winds are associated with an anomalous anticyclonic flow pattern (Appendix Figure 2, top) (Gervais et al. (2019); Romanou et al. (Under Review); Orbe et al. (Under Review)), which could contribute to the reduced heat fluxes and subsequent changes in NADW production. Therefore, to more convincingly link the surface wind speed changes to the stratospheric ozone changes aloft, we next examine results from the fixed SST experiments.

Figure 12 shows the ozone-induced zonal wind and temperature changes averaged over the last twenty years of the fixed SST and SIC experiments in which the ensemble mean ozone 4xCO_2 evolution from LINOZ is prescribed (Fig. 12 a,b). Recall that in the fixed SST experiment, only the ozone evolution differs from the preindustrial control simulation, as CO_2, SSTs and SIC are all set to preindustrial values. Comparisons with results from the fully coupled LINOZ "fast" response (see Fig. 5a,c) reveal a very similar picture. This similarity between the fully coupled fast response and the fixed SST and SIC experiment is striking, both featuring a similar change in the NH jet associated with enhanced temperature gradients in the lower stratosphere as first reported in CP2019.

Comparisons of the 850 hPa zonal winds and surface temperatures over the North Atlantic (Fig. 12c,d) also reveal a strikingly similar response between the fully coupled ensemble and the fixed SST experiments (compare with Fig. 6a,c). Note this similar response extends to sea level pressure as well (Appendix Figure 2, bottom). This result is interesting as it suggests that over the North Atlantic stratospheric ozone changes alone can result in a significant reduction in the near surface winds that is on the same order (if not larger than) the 4xCO_2 response. In our coupled atmosphere-ocean model this additionally results in heat flux changes that are large enough to reduce NADW production, resulting in a significant (i.e. 30-40%) change in AMOC strength.

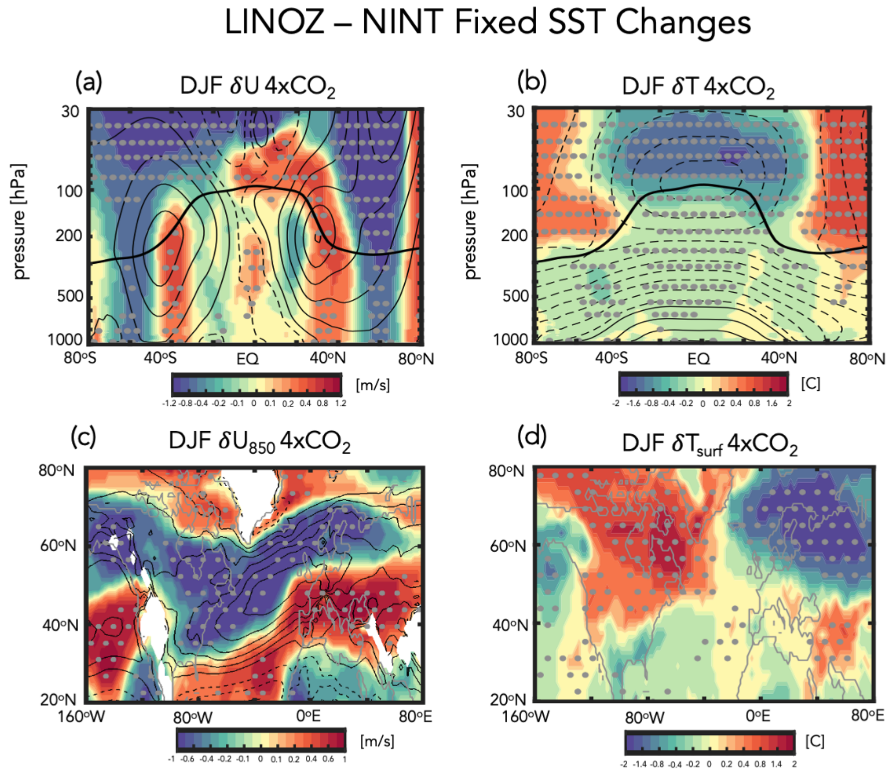

FIG. 12. Top panels: Colors show the $4xCO_2$ ensemble mean response in zonal mean zonal winds, U (a), temperatures, T (b), 850 hPa zonal winds, U_{850} (c) and surface temperature, T_{surf} (d) in the AMIP experiments in which the time-evolving $4xCO_2$ ensemble mean LINOZ ozone response is prescribed. Note that SSTs, SICs and background CO_2 are all set to preindustrial values. Averages are shown over the last 20 years (years 40-60) of the integrations. Black contours, where shown, denote climatological mean DJF values (U contour interval: 8 m/s; T contour interval: 10 C; U_{850} contour interval: 2 m/s). Stippled regions are statistically significant and the black thick line in the top panels shows the climatological mean tropopause in the preindustrial control simulation.

4. Conclusions

Here we have used the NASA GISS coupled atmosphere-ocean high-top model (E2.2-G) to examine how coupled changes in stratospheric ozone and the ocean circulation both influence the $4xCO_2$ response of the NH midlatitude jet. Our key results are as follows:

- The NH midlatitude jet response to $4xCO_2$ is modulated by coupled feedbacks from both stratospheric ozone and the AMOC, which occur of "fast" (5-20 year) and "long" (100-150 year) timescales, respectively.

- In the “fast” response, the zonal mean jet weakens (strengthens) on its poleward (equatorward) flank, consistent with reduced LS temperature gradients associated with ozone loss in the tropics. Zonally, this jet change is expressed as a negative NAO-like pattern, consisting of weaker zonal surface winds over the North Atlantic, consistent with the findings in CP2019.

- The weaker winds over the North Atlantic are associated with increased (primarily latent) heat fluxes into the ocean, which initially result in warmer SSTs over the subpolar gyre region, reducing NADW production and leading to more rapid weakening of the AMOC.

- A reduced AMOC leads to widespread cooling over the Arctic which enhance mid-to-lower tropospheric temperature gradients, resulting in a poleward shift of the NH midlatitude jet. This “long” response is consistent with previous studies showing that a weakening of the AMOC results in a stronger and poleward shifted jet in the NH (e.g., Bellomo et al. (2021); Orbe et al. (Under Review); Liu et al. (2020); Zhang et al. (Submitted)).

Taken together, the findings listed above indicate that the stratospheric ozone feedback on the NH midlatitude jet reported in CP2019 is coupled to the behavior of the AMOC during the “fast” response, wherein the jet weakens over the North Atlantic. In our model, this wind response extends to the surface, resulting in reduced heat fluxes out of the subpolar gyre region and a more rapid decline of the AMOC. On longer timescales, these changes in the AMOC subsequently drive a poleward shift in the NH midlatitude jet. Unlike the “fast” response, this “long” timescale response in the NH jet to changes in stratospheric ozone has not been previously reported, to the best of our knowledge. This may reflect the fact that many of the stratosphere resolving chemistry climate models that are used to inform future projections of stratospheric ozone (Eyring et al. (2008); Fahey et al. (2018)), are not always run coupled to an interactive ocean (Morgenstern et al. (2017)). Among those that are run coupled to a dynamic ocean, our results will, of course, need to be tested to assess robustness.

Another intriguing result from this study is that the stronger decline of the AMOC in the LINOZ ensemble does not appear to be a random occurrence. Rather, in our model, the “fast” ozone and “long” AMOC feedbacks on the NH jet are coupled through surface-wind driven changes in heat fluxes into the ocean. Key here is the fact that this sensitivity in the AMOC is driven only by changes in stratospheric ozone, which we have isolated from changes in other trace gases and aerosols. Thus, while previous studies (Rind et al. (2018)) have identified an important influence of interactive composition on the AMOC, they have mainly implicated the indirect effect of aerosols on clouds through changes in sea surface temperatures and how these impact P-E (and net surface freshwater forcing). To the best of our knowledge, no study has previously demonstrated an impact of stratospheric ozone changes alone on the AMOC response to a quadrupling of CO_2. Despite the different mechanisms at play, however, are results are generally consistent with those from Rind et al. (2018) in highlighting the need for renewed focus on surface flux observations to help assess overturning stability.

An important caveat with our results is related to known biases in vertical mixing and NADW production in the ocean component of the GISS model (Miller et al. (2021); Romanou et al. (Under Review)) which likely explain why the low-top version of the coupled atmosphere-ocean climate model (E2.1-G) exhibits a more sensitive AMOC response to a quadrupling of CO_2, compared to some other models (Bellomo et al. (2021)). At the same time, the high-top model employed in this study is much less sensitive, as the AMOC weakens by ~10 SV in response to 4xCO_2, compared to a complete collapse in E2.1-G (see Figure 31 in Rind et al. (2020)). That study showed that this may be related to differences in the parameterization of rainfall evaporation associated with moist convective precipitation, which they show has a strong influence on the AMOC sensitivity in the GISS model via its effect on moisture loading in the atmosphere. While an exhaustive comparison between the models is beyond the scope of this study, the relevant point here is that the 4xCO_2 AMOC response simulated in the E2.2-G NINT ensemble is well within the CMIP5 and CMIP6 ranges documented in Mitevski et al. (2021) (see their Supplementary Figure S3).

Finally, our results linking the fast timescale jet response to the ensuing AMOC changes underscore the profound impact that changes in lower stratospheric winds alone can have on surface climate, as highlighted in Sigmond and Scinocca (2010). Quite remarkably, our fixed SST and SIC experiment showed that these lower stratospheric wind changes are driven primarily by changes in ozone and not by background changes in CO_2 or in sea surface boundary conditions. Taken together, our results suggest that more attention needs to be paid to understanding the time-evolving response of the coupled Earth system to future ozone changes, with a focus on changes in ocean heat transport and how these feed back on the NH jet stream.

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Data availability statement. The OMA and NINT GISS E2.2-G data used in the study are available at the CMIP6 archive via the Earth System Grid Federation (`https://esgf-node.llnl.gov/`). The specific simulations used here are the piControl, abrupt-2xCO2, and abrupt-4xCO2 r1i1p1f1 (NINT) and r1i1p3f1 (OMA) runs. Output from the additional three NINT 4xCO2 simulations as well the four-member LINOZ ensemble is available at `https://gmao.gsfc.nasa.gov/gmaoftp/corbe/AMOC_Linoz/`. All GISS ModelE components are open source and available at http://www.giss.nasa.gov/tools/modelE/.

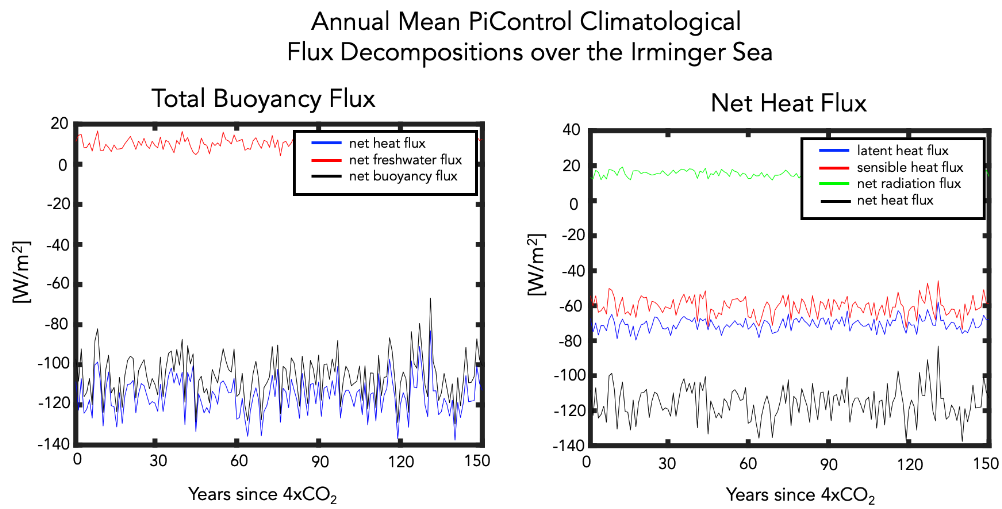

FIG. A1. Left: Decomposition of the net surface buoyancy flux (black) into its contributions from net heat (blue) and net freshwater (red) fluxes. Right: Further decomposition of the net surface heat flux (black) into contributions from latent heat fluxes (Q_E (blue)), sensible heat fluxes (Q_H (red)), and combined solar and longwave radiative fluxes (Q_S+Q_L (green)). Results are shown for 150 years of the NINT preindustrial control (PiControl) simulation, evaluated over the Irminger Sea.

APPENDIX

A1. Appendix Figures

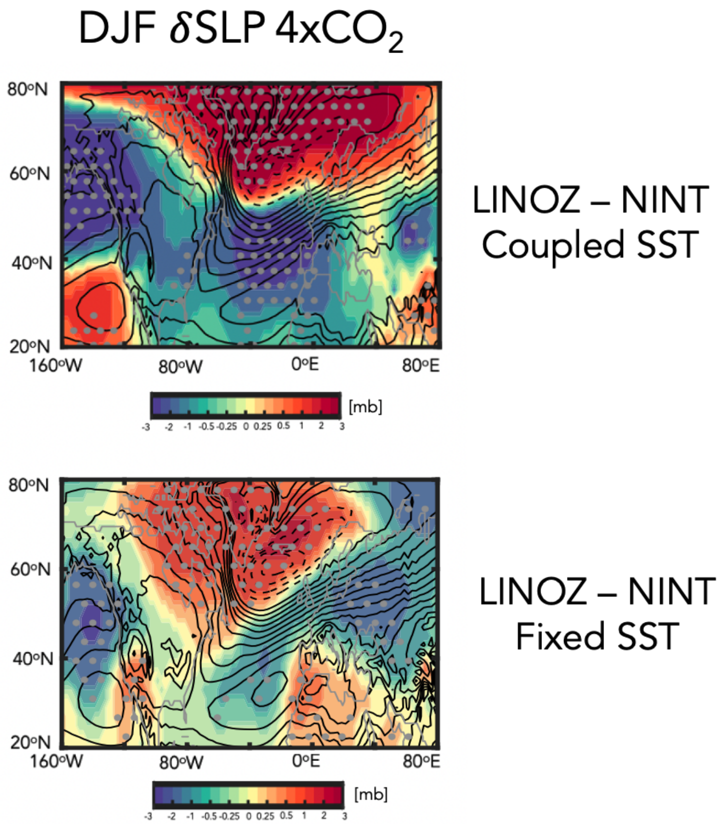

FIG. A2. Top panel: Colors show the LINOZ minus NINT ensemble mean difference in the December-January-February (DJF) "fast" response of the sea level pressure to an abrupt quadrupling of CO_2. Results are shown for the fully coupled atmosphere-ocean simulations. Bottom panel: The ensemble mean response in sea level pressure in the AMIP experiments in which the time-evolving 4xCO_2 ensemble mean LINOZ ozone response is prescribed. Note that SSTs, SICs and background CO_2 are set to preindustrial values. Black contours denote climatological mean DJF values (contour interval: 10 mb). Stippled regions are statistically significant.

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